State of California Resources Agency Department of Water Resources Southern District Resources Assessment Branch Groundwater Section

Technical Information Record SD-07-03

A Summary of the Drilling and Completion of the Borrego Water District Monitoring Well MW-4 with Geologic and Geophysical Observations and Initial Interpretations

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This TIR is primarily a working paper and is subject to revision or replacement

Introduction

Monitoring well MW-4 was completed as part of a project partially funded by a Local Groundwater Management Assistance Act¹ (AB303) grant awarded to the Borrego Water District to help develop a monitoring well network in the Borrego Valley groundwater basin. The network is being established to increase the district's ability to monitor and record the groundwater levels and quality in the basin. The enhanced data collection will help the Borrego Water District optimize its groundwater management plan to mitigate a significant overdraft condition in the basin. Other monitoring wells in the network, MW-1 and MW-3, were completed in 2004, the latter being documented in Ross, 2005. Monitoring wells MW-4 and MW-5 were completed in 2006 and are documented here and in Ellis and Ross, 2007, respectively. In addition to these recently drilled monitoring wells, several existing public and private supply wells and private irrigation wells, both active and abandoned, have been included as part of the network. An existing network of observation wells owned by San Diego County and monitored by the Borrego Water District provides additional groundwater level and quality data.

MW-4 is a single completion well located inside the southern perimeter fence near the east end of the runway at the Borrego Valley Airport on Palm Canyon Drive (Figure 1). The airport site for MW-4 was selected to provide groundwater level and quality data representative of the immediate area and to test the aquifer specifically for its suitability to supply a municipal water well. Geologic and geophysical data obtained at this site also provides a means to delineate the thicknesses and identities of underlying formations and examine their relation to the regional aquifer system.

The site surface is mantled in aeolian dune sand with a moderate brush overgrowth. The drill location was leveled, and a working pad was built up with the sand to accommodate the drill rig, auxiliary vehicles, and other equipment. The drilling was completed by SoCal Pump and Well Drilling, Inc. of Riverside, California using the direct rotary technique with a bentonite-based drilling mud. Steel mill-tooth tricone bits were

¹ California Water Code Section 10795 et seg.

used for drilling the borehole, and a Boart Longyear diamond coring bit with a 10-foot core recovery barrel was used for coring of selected sections.

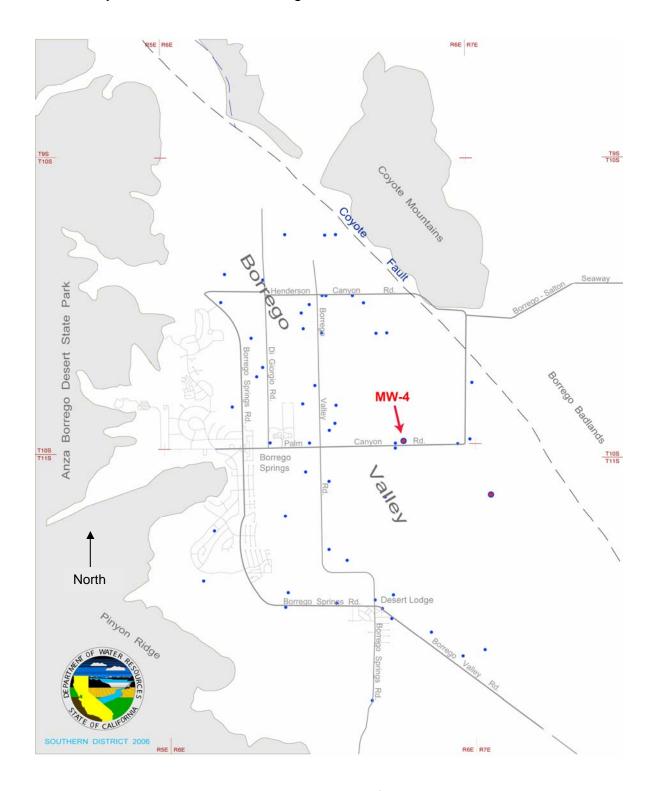


Figure 1 - Location of MW-4

During the drilling, cutting samples were collected from the returning mud at the conductor casing and, occasionally, at the shaker table as a supplement. Samples were collected using a steel-screen strainer and washed in water to remove residual drilling mud for classification and storage. Recovered core samples were cleaned, photographed, and described. The core samples are stored at the Borrego Water District office in Borrego Springs.

Well boring and completion summary

Drilling at the airport site for MW-4 commenced on July 6, 2006, at 1700 and concluded at approximately 1200 on July 26, 2006. The uppermost part of the borehole was drilled to 24 feet below land surface (bls) using a 17-inch steel mill-tooth tricone bit, reamed, and completed with a 24-foot length of 16-inch diameter steel conductor casing set without external grout. A sample of the surface aeolian sand was collected adjacent to the borehole to serve as a reference for identifying contamination in cutting samples. Samples were collected from mud returns at the borehole from depths of 10 and 20 feet bls. These first two samples were examined and found to contain a significant surface sand fraction, which likely was caused by caving or similar contamination in the absence of the conductor casing. Samples obtained at 30 feet bls, after the conductor was set and the borehole was cleared, appeared more representative of the formation encountered. Drilling continued to the total depth of the borehole (927 feet bls) using a 9-inch steel mill-tooth tricone bit. Selected core sampling intervals were drilled with the coring bit and reamed using the 9-inch tricone bit.

Sampling intervals were gauged to occur at 10-foot increments by estimating the depth from the top swivel position on the drill rig's derrick. This estimate did not account for the progressive delay or "bottoms up" time in sample collection as the borehole deepened. Initial delays were negligible and the samples collected were considered true representations of the lithology at the estimated depths. The delay at 650 feet bls, however, was measured at approximately 15 minutes, which corresponds to an offset of approximately 10 to 20 feet between the estimated depth and the sample depth of the

recovered sample at the deeper intervals. This offset is taken into consideration when comparing the lithologic log with the down-hole geophysical log results.

Coring was attempted at four 10-foot intervals using a Boart Longyear coring bit and 10-foot core barrel. The first attempt was at 105 to 115 feet bls, however, no sample was recovered, presumably because of insufficient consolidation in the formation. Partial core samples were recovered from 847 to 857 feet bls (6-foot length), 867 to 877 feet bls (0.7-foot length), and 887 to 897 feet bls (0.46-foot length with loose cobbles).

A geophysical survey of the entire 927-foot deep borehole was completed by Pacific Surveys on July 27, 2007. The survey included measurements of the spontaneous potential (SP), natural gamma-ray emission, resistivity (long and short normal, single point, laterolog), sonic velocity, variable density, and borehole dimension using an x-y caliper. The data from the surveys are included as Appendix 1.

After the survey was completed, the borehole was grouted with bentonite from the bottom (927 feet bls) to 395 feet bls. The well was completed in the remaining open borehole to 390 feet bls. The completion consists of the 16-inch diameter conductor to 24 feet bls, a 6-inch diameter casing from the ground surface to 85 feet bls, and 6-inch diameter screens from 85 to 390 feet bls. Final surface completion was made flush with a recessed vault. The static depth to water in the completed well on August 4, 2006, was 94 feet bls.

Step drawdown and constant rate pumping tests were conducted on the completed well by Borrego Water District personnel. Results from these tests indicate that the well can sustain a 250-gallon per minute (GPM) yield under the present groundwater and well conditions. Water samples were drawn by Borrego Water District personnel and sent to E.S. Babcock & Sons laboratories for analysis. The major ion analysis results show that the water is a sodium sulfate type with a total dissolved solid (TDS) concentration of 720 milligrams per liter (mg/l). Saturation indices were calculated from the water quality analysis results using the U.S. Geological Survey PHREEQC

geochemical modeling software (Parkhurst and Appelo, 1999). The sample was found to be saturated with respect to calcium carbonate, consistent with the light cement coating found on some sediment grains in the lithologic samples. Table 1 shows the major ion chemistry and saturation indices.

Table 1 – Water Quality Analysis and Saturation Indices for MW-4, January 8, 2007

Water C	Quality A	nalysis								
Ca	Mg	Na	K	HCO ₃	SO_4	CI	NO_3	F	TDS	рН
78.0	15.0	130.0	9.4	110.0	330.0	66.0	2.4	0.5	720.0	7.8
Concen	trations	in mg/l								
Saturat	ion Indic	es								
Calc	cite (CaC	CO ₃)		Gypsum	n (CaSO₄	•2H ₂ O)				
	0.27				-0.69					

Borehole description

The overall geology of the units encountered during the drilling ranges from coarse-grained sand and gravel to clay sedimentary deposits consistent with the alluvial, lacustrine, and deltaic deposits described regionally (e.g., Moyle, 1982; Cassiliano, 2002). Thin interbedding of sandy to clayey sediments are found to varying degrees throughout the borehole depth as indicated by the typically erratic trace of the long normal resistivity measurements interpreted in many places as reversals. The thin interbedding is estimated to occur on a scale of 5 or less feet and overprints a broader cyclic variation between sandy and clayey layers that occurs on a scale of tens of feet. The thinner interbedding is not continuously resolved in the lithologic log, presumably because of vertical mixing in the returning mud stream and the relatively fast drilling rate.

Generally, an overall decrease in the porosity with increasing depth is suggested by a steady increase in the sonic velocity as represented by a decrease in the return time

(delta T) values. This overall decrease in porosity can be generally attributed to increasing sediment compaction and limited induration with increasing depth, although core samples from more than 800 feet bls suggest that consolidation and induration has not occurred to a significant degree. Variations in sonic velocity within the overall increasing trend are noted to coincide with lithologic changes in the strata from clay- to sand-dominated layers.

Sand-dominated layers are prevalent from approximately 140 to 330 feet bls, 480 to 780 feet bls, and more than 820 feet bls as interpreted from the geophysical logs. These interpretations generally agree with descriptions of poorly-graded sand (SP) and clayey sand (SC) recorded in the lithologic log. Some portions of these intervals, however, are generally described as clay dominated (CL), though the detailed descriptions in the lithologic log include a significant sand component.

The sand-dominated layers are typified by relatively low but varied gamma emission. An exception is found from approximately 240 to 285 feet bls where a spike in gamma emission occurs along with a number of peripheral increases in gamma emission. The spike in gamma emission may be associated with an unconformity. Clay-dominated layers are typified by a relatively high and varied gamma emission. A second possible unconformity, also represented by a spike in gamma emission, is located at approximately 400 feet bls within clay-dominated layers.

Short and long normal resistivity traces generally coincide in sand-dominant layers, but separate in a number of intervals with the long normal resistivity values decreasing relative to the short normal resistivity values. These separations are most notable from 300 to 340 feet bls and 820 to 885 feet bls. The separation suggests that the formation water has a relatively high TDS concentration. The separation is more pronounced at depths greater than 820 feet bls, which suggests that the TDS concentration of the formation water increases with greater depth representing either vertical isolation between aquifers or simple stratification within a single common aquifer.

Large fluctuations affecting both long and short normal resistivity traces occur throughout the depth of the borehole. These fluctuations often coincide with variations in the borehole diameter as recorded by the caliper log and are thus considered exaggerated relative to the true resistivity measures of the formations and mud-filled borehole.

Separations between long and short normal resistivity traces, where the short normal resistivity values decrease relative to the long normal resistivity values, coincide with borehole diameter increases in clay-dominated layers. The relative decrease in short normal resistivity values is attributed to an increase in the influence of the low resistivity drilling mud. Infiltration of drilling mud into the borehole wall is also a possible contributor to the observed separation, though the occurrence of this in clay-dominated layers is likely minimal.

Core sampling

Core sampling was completed at four depth intervals. A sample was attempted from 105 to 115 feet bls with no recovery. The interval was anticipated to coincide with the water table; however, an apparent lack of consolidation of the formation prevented recovery of any sample. Core recovery was also attempted between 847 and 857 feet bls, 867 and 877 feet bls, and 887 and 897 feet bls.

Between 847 and 857 feet bls, a 6.4-foot core sample was recovered in three pieces: a 6-foot segment, a 0.25-foot segment, and a 0.15-foot segment. The core composition is uniform throughout the sample lengths of all three segments. The grain sizes are dominantly fine sand and silt (<1 mm), with lesser coarse sand and fine gravel (up to 5 mm). The sample matrix is dominantly grain supported with some clay present in the interstices. No cementation is noted within the matrix (negative HCl reaction). Coarse sand and gravel are primarily colorless to milky quartz in angular- to subangular-grains. Coarse gravel (20 to 30 mm) is present in isolated occurrences.

Between 867 and 877 feet bls, a 0.7-foot core sample was recovered. The composition is identical to the core sample collected between 847 and 857 feet bls. The lack of recovery of more core material from this interval suggests that the formation is not well consolidated at this depth.

Between 887 and 897 feet bls, a final 0.46-foot core sample was recovered. The core sample also has an identical composition to the core samples noted above. The core sample was retained in the base of the barrel. Also retained in the barrel on top of the core sample were several coarse pebble- and cobble-sized rock fragments lying loose. The rock fragments are gneissic, plutonic, and sandstone pebbles and small cobbles that are well rounded and, in the case of the larger pieces, cut in a radius by the coring bit suggesting that they were part of a larger boulder in place. Also included are several dark-grey, fossiliferous claystone clasts that are rounded and slightly weathered.

Water-bearing zones

Based on the lithologic log and the geophysical data, three apparently distinct water-bearing zones are identified. The upper zone is between the water table at 94 feet bls and approximately 350 feet bls. The middle zone is between 430 and 780 feet bls. The lower zone extends from approximately 820 feet bls to, at least, the bottom of the borehole at 927 feet bls.

The interval separating the upper and middle zones has markedly increased clay content as indicated in the lithologic log and corroborated by an increase in gamma emission. The borehole diameter varies significantly in this interval, possibly caused by lateral "walking" of the drill bit in response to an increase in the clay content of the sediments. Resistivity values decrease in this interval, with separations between long and short normal traces indicating a greater influence of less resistant drilling mud in the short normal trace in response to increased borehole diameter.

The interval separating the middle and lower zones also has an apparent increase in clay content. This intervening layer is identified by an increase in the gamma emission and notable slowing in the sonic velocity. Resistivity values drop sharply across this interval as well; however, a significant variance in the borehole diameter as shown by the caliper log is likely influencing these data.

Between 820 and 880 feet bls, a significant drop in overall clay content relative to the middle zone is indicated by a drop in gamma emission and an increase in sonic velocity. The core samples recovered within this interval were found to have a grain-supported matrix with lesser clay in the interstices and low overall consolidation, consistent with the characteristics indicated by the geophysical logs.

A small spike in the gamma emission at approximately 888 feet bls roughly coincides with an increase in overall resistivity and follows an abrupt increase in the borehole diameter between 882 and 888 feet bls. The core recovered from the deepest interval (887 to 897 feet bls) includes the small cobbles and pebbles not seen in shallower cores. These observations are not reflected in the lithologic log possibly because of the subtlety of the change and vertical mixing within the borehole. The abrupt changes noted, coupled with the recovery of the cobbles and pebbles in the deepest core, suggest that the borehole encountered a formation or facies change at approximately 880 feet bls, with the gamma spike at 888 feet bls marking a possible unconformity.

Conclusions

Future analysis of the collected lithologic samples and geophysical data will be carried out to correlate the strata encountered in the boring of MW-4 with the regional geology. Preliminary estimates of the identity of the various formations encountered suggest that most of the borehole intercepts older alluvium and formations of the Palm Spring Group as described by Cassilliano (2002). The recovery of pebbles, cobbles, and rounded fossiliferous-claystone between 887 and 897 feet bls, and the possible

existence of an unconformity at 888 feet bls indicated by the gamma emission spike suggest the occurrence of a formation or facies change, possibly coinciding with the basal part of the Palm Spring Group.

References

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Appendix 1

Borehole geophysical survey results

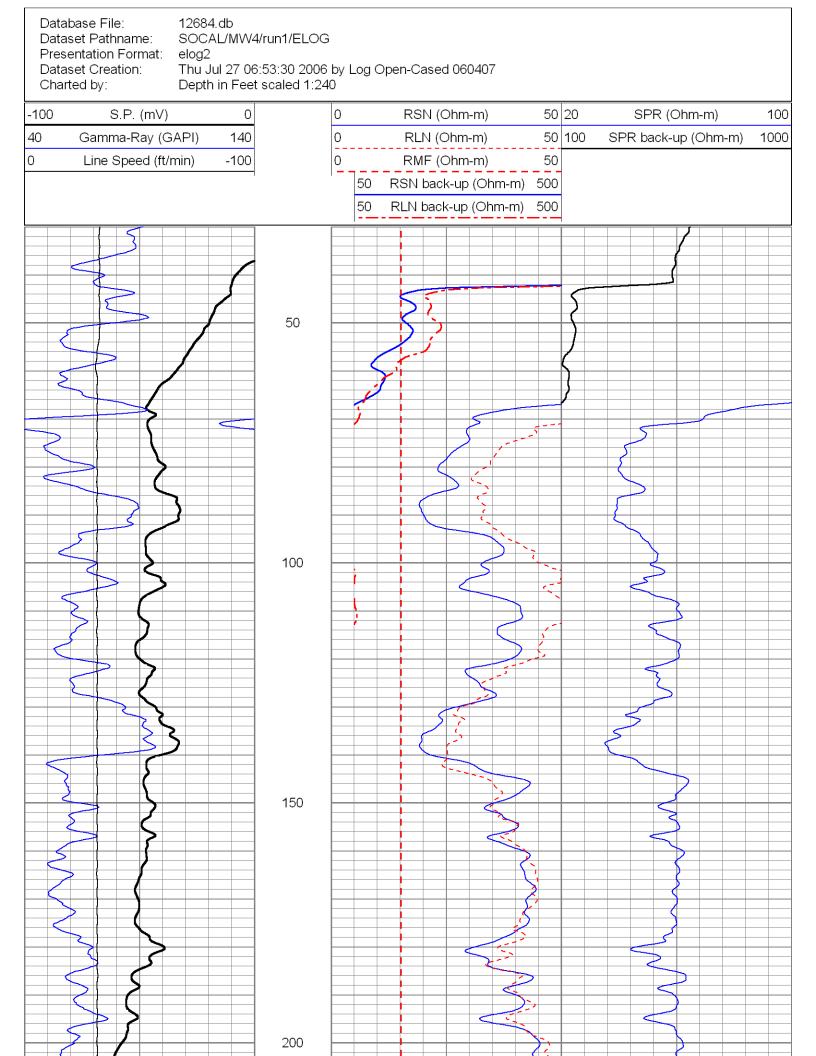
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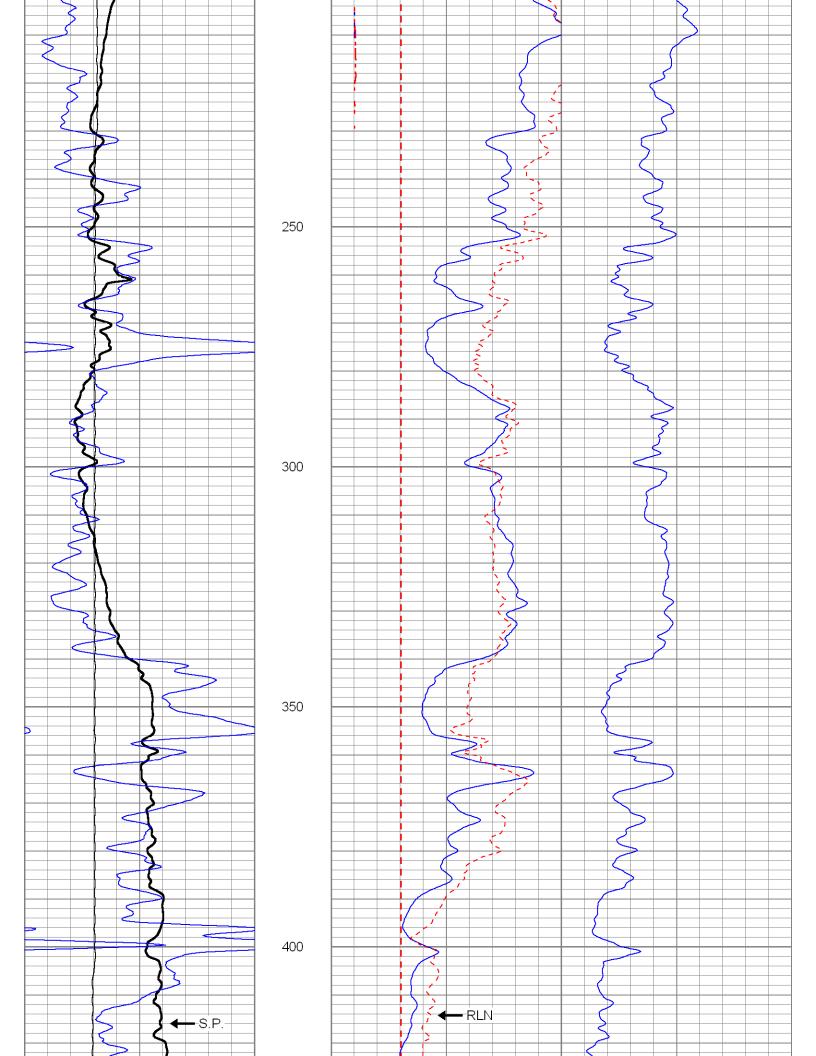
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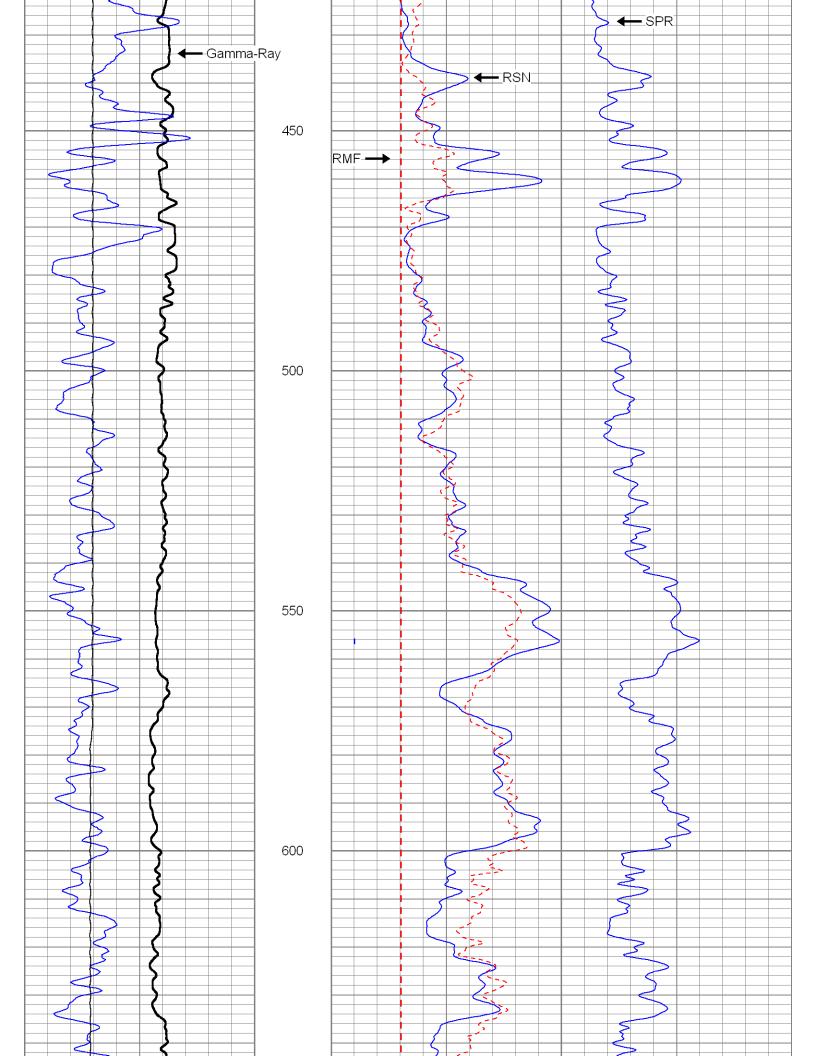
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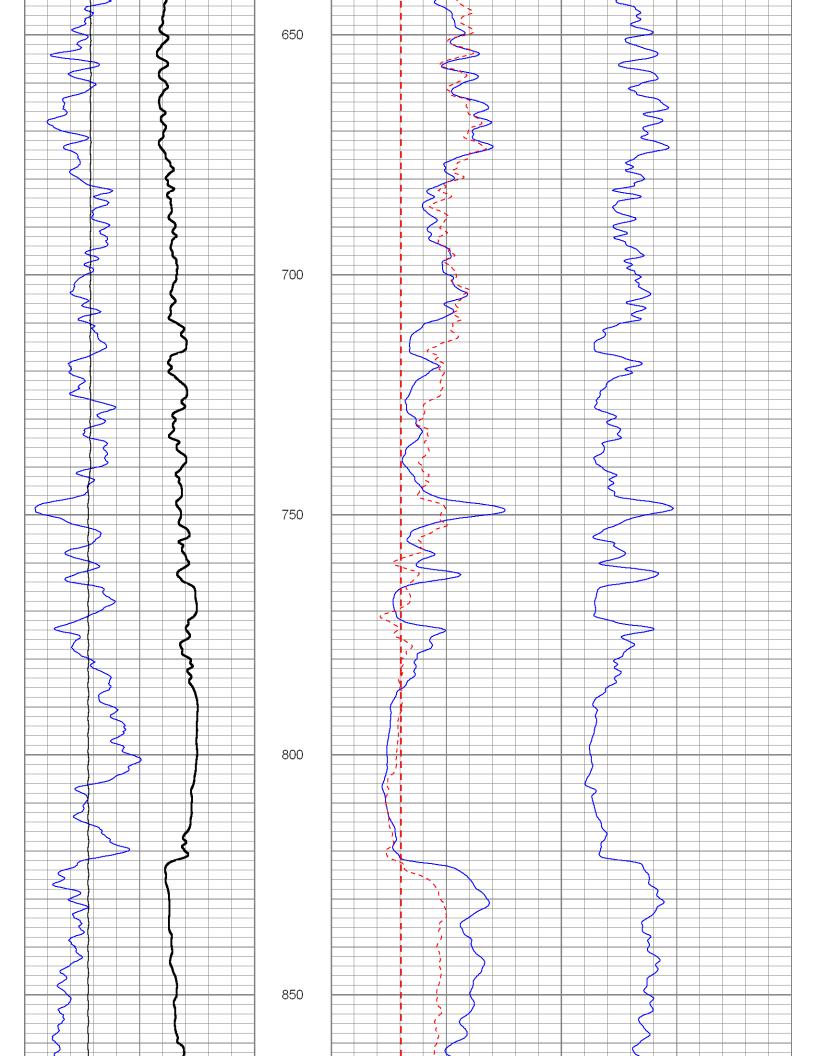
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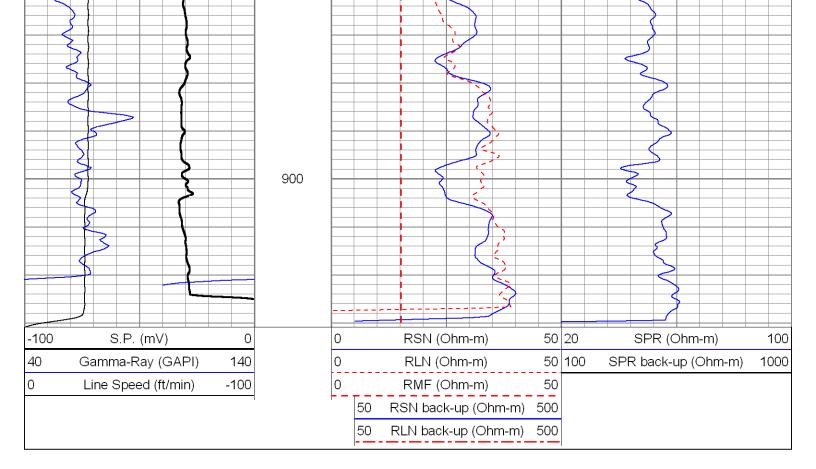
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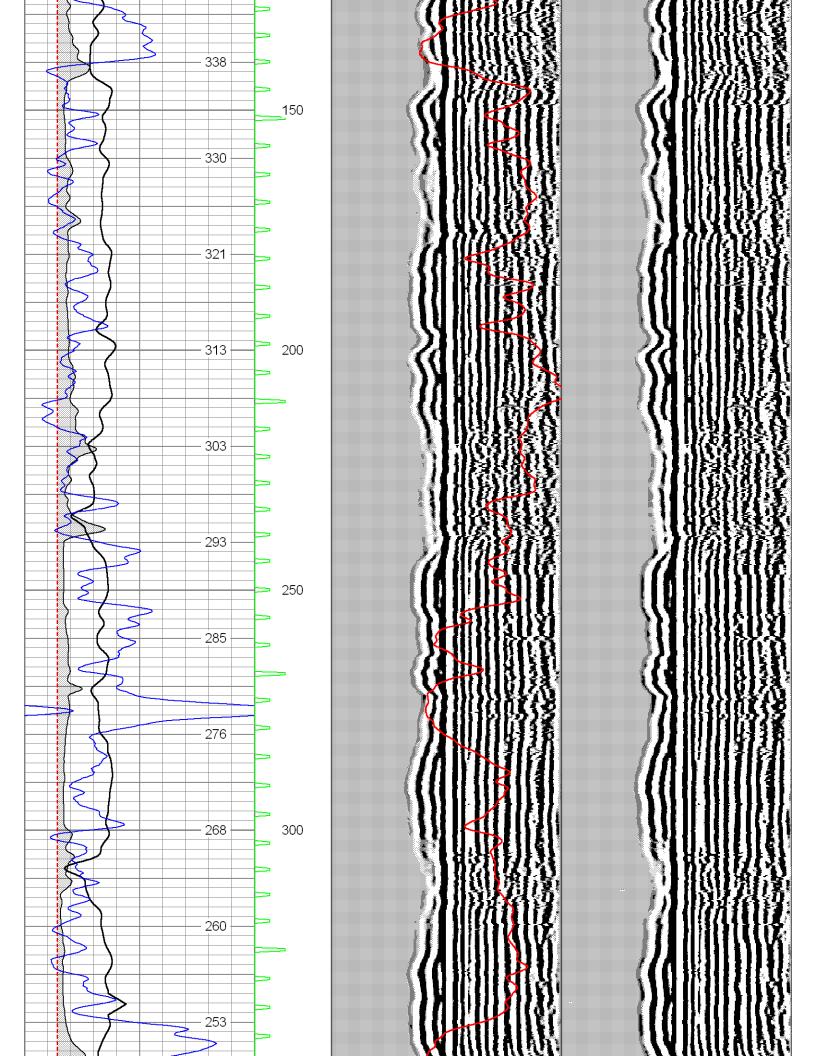
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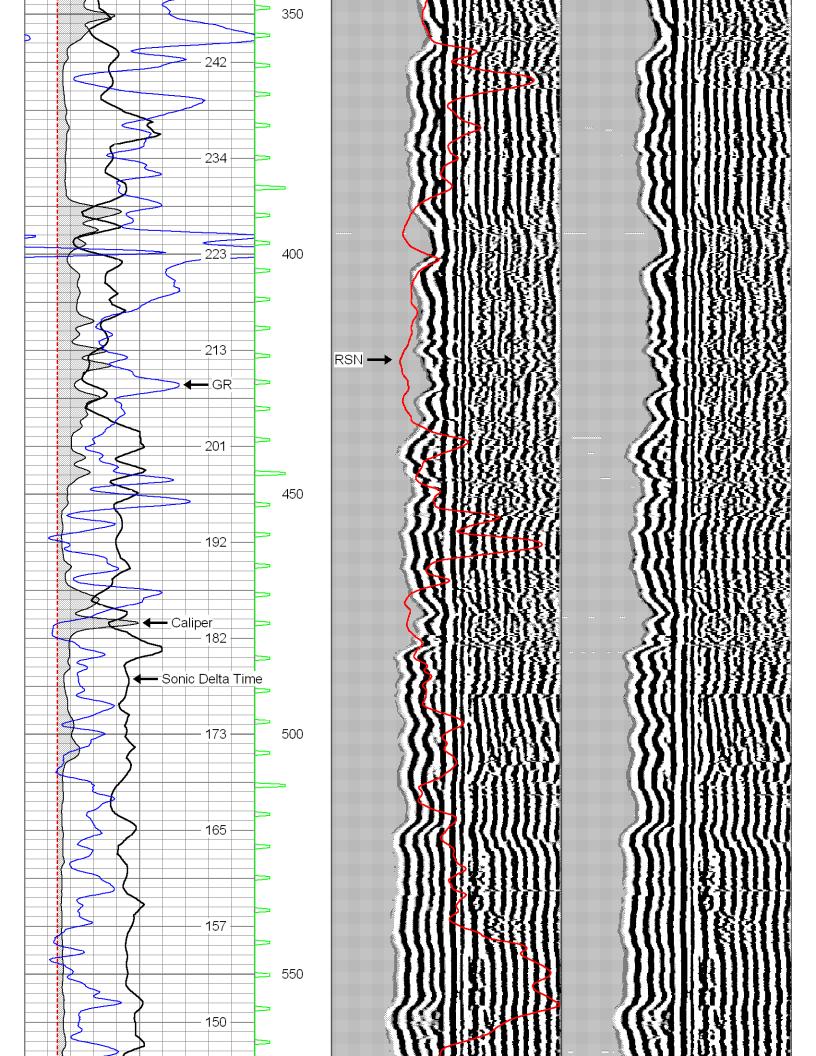
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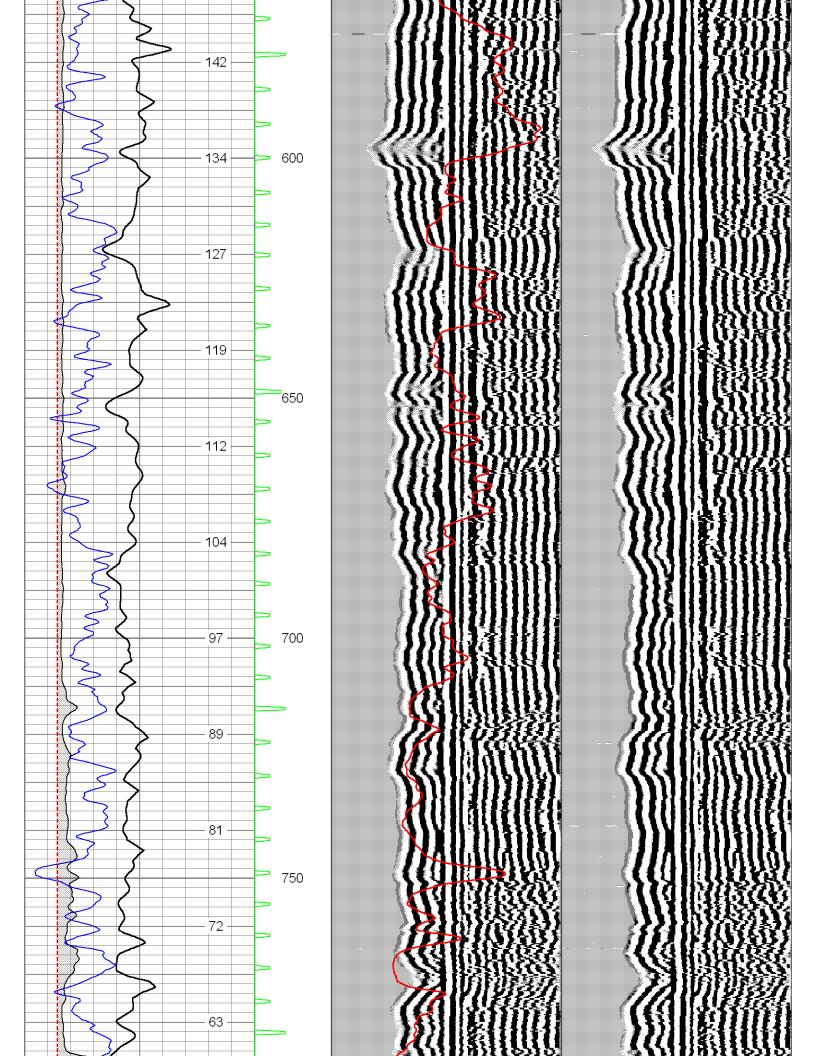
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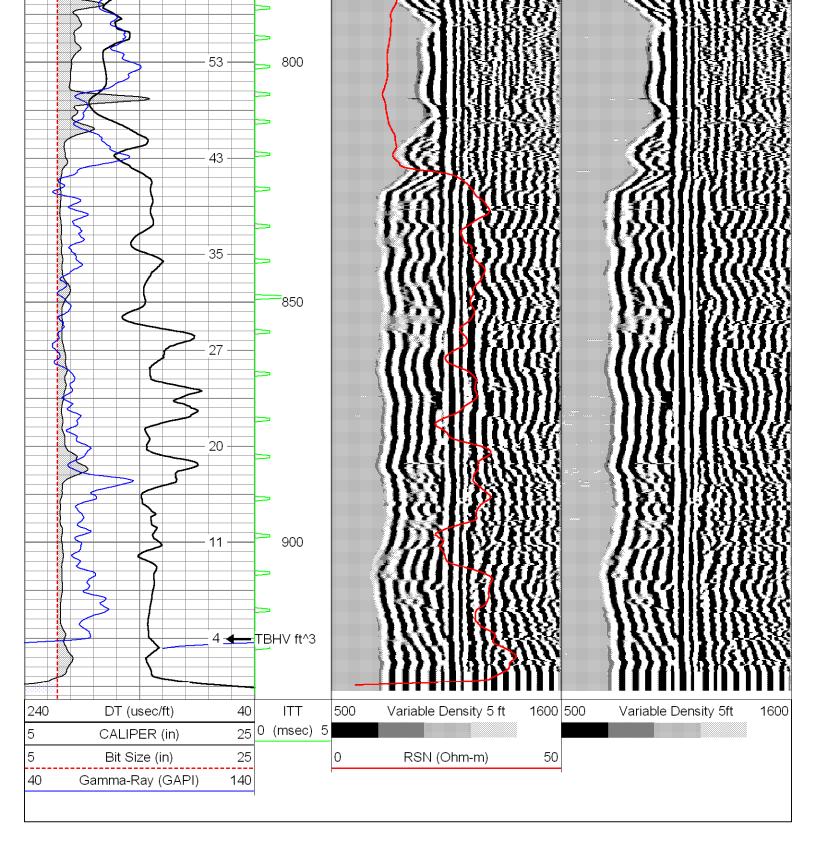
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Dataset Pathname: SOCAL/MW4/run1/LL3
Dataset Creation: Thu Jul 27 07:41:47 2006 by Log Open-Cased 060407

12 Serial Number: GROH Tool Model: Performed: Thu Feb 12 09:06:37 2004 Calibrator Value: 162.0 GAPI Background Reading: 46.0 Calibrator Reading: 174.5 Sensitivity: 1.2612 GAPI/ RLL3 (Resistivity Laterolog 3) Calibration Report: Serial Number: 294 Tool Model: M&W Performed: Mon Jul 25 11:54:08 2005 System Reading Calibration Reference 2.500 0.029 Ohm-m 0.042 5.000 0.410 50.000 2.079 250.000 4.401 500.000 Database File: 12684.db Dataset Pathname: SOCAL/MW4/run1/LL3 Presentation Format: Dataset Creation: Thu Jul 27 07:41:47 2006 by Log Open-Cased 060407 Charted by: Depth in Feet scaled 1:240 0 Line Speed (ft/min) -100 RLL3 (Ohm-m) 50 50 140 0 40 Gamma Ray (GAPI) RSN (Ohm-m) 0 RLN (Ohm-m) 50 0 50 RMF (Ohm-m) CLL3 250 0 500 50 RLL3 back-up (Ohm-m) 50 500 RSN back-up (Ohm-m) 50 RLN back-up (Ohm-m) 500

